

European ammonoid diversity questions the spreading of anoxia as primary cause for the Cenomanian/Turonian (Late Cretaceous) mass extinction

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ABSTRACT

Ammonoid diversity patterns show that the spreading of oceanic anoxia is not the initial and major kill mechanism for the Cenomanian/Turonian mass extinction as usually suggested. In the Anglo-Paris Basin and the Vocontian Basin, the drop of ammonoid species richness starts around the middle/late Cenomanian boundary, i.e. 0.75 myr before the occurrence of anoxic deep-water sediments. The stepwise extinction of first heteromorphs and then acanthoceratids is incompatible with the rise of the oxygen minimum zone. Moreover, shelf environments of these basins remained well oxygenated during the Cenomanian/Turonian boundary interval. Thus, we stress that other causative mechanisms initiated the ammonoid extinction even if anoxia subsequently participated in the demise of marine ecosystems.

RESUME

L'extinction de masse de la limite Cénomannien/Turonien (Crétacé supérieur) est habituellement considérée comme un exemple typique d'extinction globale causée par la remontée d'eaux anoxiques. Cependant, l'analyse des changements de biodiversité des ammonoïdés durant l'intervalle Cénomannien-Turonien inférieur démontre que cet événement océanique anoxique (OAE2) n'est pas la cause initiale et majeure de l'extinction des ammonites. Les bassins européens (Bassin anglo-parisien et Bassin vocontien) montrent que la chute de diversité (en terme de richesse spécifique) des ammonites s'amorce autour de la limite Cénomannien moyen/supérieur. Le déclin des ammonites se produit donc 0,75 million d'années avant le développement des eaux anoxiques. A ce décalage temporel entre les ammonites et l'anoxie s'ajoute une incompatibilité entre l'expansion d'eaux anoxiques vers les plateformes et le déclin prématuré des hétéromorphes (plutôt pélagiques) par rapport aux acanthocératidés (plutôt necto-benthiques). De plus, l'enregistrement stratigraphique des domaines peu profonds de ces bassins montre clairement que les eaux sont toujours largement oxygénées pendant la limite Cénomannien/Turonien. Il est donc probable que des mécanismes autres que l'anoxie doivent avoir initié l'extinction de masse de la limite Cénomannien/Turonien (au moins en ce qui concerne les ammonites), même si l'anoxie a ensuite partiellement participé au bouleversement des écosystèmes marins à cette période.

Introduction

The Cenomanian-Turonian is one of the most studied marine stratigraphic intervals in the Cretaceous. This focus is prompted by the occurrence of peculiar biotic and abiotic events, such as a moderate mass extinction (Raup & Sepkoski 1984, 1986), the highest sea level of the Mesozoic (Haq et al. 1987), some of the highest atmospheric CO₂ concentrations (Berner 1994; Bice & Norris 2002), a temperature peak (Jenkyns et al. 1994), a global oceanic anoxic event (Arthur et al. 1987; Schlanger et al. 1987), a δ¹³C positive excursion caused by a major perturbation of the global carbon budget (Scholle & Arthur 1980), and the deposition of organic-rich sediments (Herbin et al.

1986; Schlanger et al. 1987), among others. Therefore, the Cenomanian-Turonian is a key interval to analyze the interactions between biotic and abiotic events.

The Cenomanian/Turonian boundary (CTB) is characterized by a worldwide diversity drop of a number of groups, such as benthonic and planktonic foraminifers, calcareous nannoplankton, scleractinian corals, irregular echinoids, aragonitic rudist bivalves, and ammonoids (Hallam & Wignall 1997). With the extinction of 26% of marine animal genera, Raup & Sepkoski (1986) recognized this event as one of the eight major mass extinctions of the Phanerozoic. Harries & Little (1999) reported the extinction of 79% of the macro-invertebrate species in the Western Interior Basin (USA).

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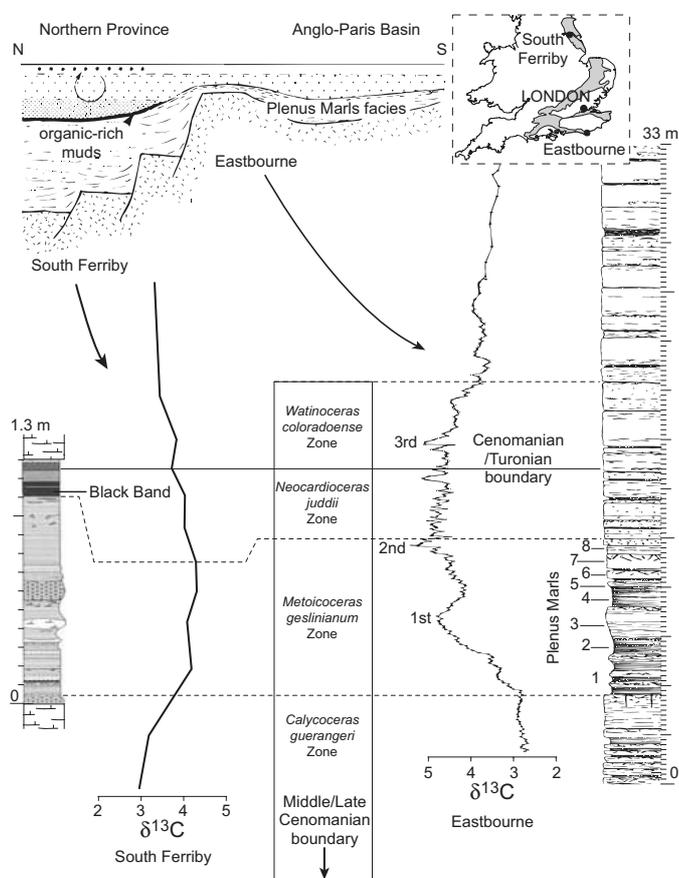


Fig. 1. Carbon isotope stratigraphy and ammonoid biostratigraphy of the Eastbourne and South Ferriby sections in the Anglo-Paris Basin. $\delta^{13}\text{C}$ curve of Eastbourne after Paul et al. (1999). $\delta^{13}\text{C}$ curve of South Ferriby after Wood & Mortimore (1995).

Within the Plenus Marls (late Cenomanian) of the Anglo-Paris Basin (north-west Europe), Jarvis et al. (1988) and Leary et al. (1989) documented the sequential decline and removal of microfaunal organisms (benthonic and planktonic foraminifers, ostracods, dinoflagellate cysts, calcareous nannofossils) in an order that corresponds to their supposed dwelling depth in the water column, thus producing a 'stepwise mass extinction' (Kauffman 1984; Hart & Leary 1991). This depth gradient is corroborated by the present-day morphology of planktonic foraminifers (Hart 1980; Bé 1982; Hemleben et al. 1989) and by oxygen isotope studies (Corfield et al. 1990; Fisher & Arthur 2002; Price & Hart 2002). Hence, several authors (e.g. Kuhnt et al. 1986; Jarvis et al. 1988; Leary et al. 1989) interpreted this progressive depth-dependent extinction of the microfaunas as the effects of a rise of the Oxygen Minimum Zone (OMZ), i.e. the spreading of hypoxic/anoxic waters onto the platforms.

A significant positive excursion in the carbon isotope record (Scholle & Arthur 1980) is also well documented during the late Cenomanian. During this interval, organic-rich sedi-

ments were deposited in the deep parts of numerous basins and oceans (Schlanger et al. 1987). These black shales, which are widely regarded as the result of increased preservation of organic matter, characterize the so-called Oceanic Anoxic Event 2 (OAE2). Hence, anoxia is generally considered the major cause of the Cenomanian/Turonian boundary (CTB) mass extinction and the killing-mechanism for the extinction of numerous organisms (e.g. rudists, Philip & Airaud-Crumière 1991; Steuber & Löser 2000; foraminifers, Groshény & Malatre 1997; ostracods, Babinot et al. 1998; radiolarians, Erbacher & Thurow 1997; O'Dogherty & Guex 2002; ammonites, Batt 1993). Therefore, the CTB mass extinction is considered a typical example of a global extinction and turnover caused by the development of anoxic waters (Kauffman & Hart 1995).

However, Gale et al. (2000) and Smith et al. (2001) have questioned the existence of this mass extinction. They argued that the faunal turnover might essentially result from a taphonomic-preservation bias generated by the spread of hemipelagic depositional environments onto the shelves resulting from an exceptionally high sea level. Thus, turnover of microfaunas during the CTB may be related to local environmental changes and the turnovers do not necessarily record increased global extinction events. Banerjee & Boyajian (1996) also documented the extinction of only 17% of foraminifer genera (within a binned-stage database), with a significant preferential bias toward discoid-agglutinated forms. The latter are usually known as more oxic-tolerant, therefore suggesting that anoxia may not have been the cause of the CTB foraminifer extinction. Finally, in contrast to all these interpretations, Monnet et al. (2003) documented a diversity drop of ammonoids in north-west European basins, thus supporting the existence of an extinction event. However, this drop in diversity seems to have predated the spread of anoxic waters and appears to contradict the hypothesis of Elder (1989) and Batt (1993), who argued that the anoxic event caused a stepwise demise of ammonoids. Even if the drop of ammonoid diversity documented by Monnet et al. (2003) confirms the existence of a moderate mass extinction, it casts serious doubt on the widely quoted anoxia as a kill-mechanism for all marine clades during the late Cenomanian. Here we discuss the timing of anoxia in relation to biodiversity patterns of ammonoids of the Anglo-Paris Basin, for which the model of anoxia-based extinction was initially proposed, and of the Vocontian Basin, for which black shales are well documented.

Timing of anoxia

During the Late Cretaceous, the Anglo-Paris Basin belonged to the north-European chalk sea. Conspicuous lateral lithologic changes led to the recognition of two major depositional provinces (Wood & Smith 1978): the northern province (eastern England, North Sea and northern Germany) and the Anglo-Paris province (southern England and Paris Basin). Within these two provinces the CTB succession is differently expressed (Fig. 1): in the northern province the succession

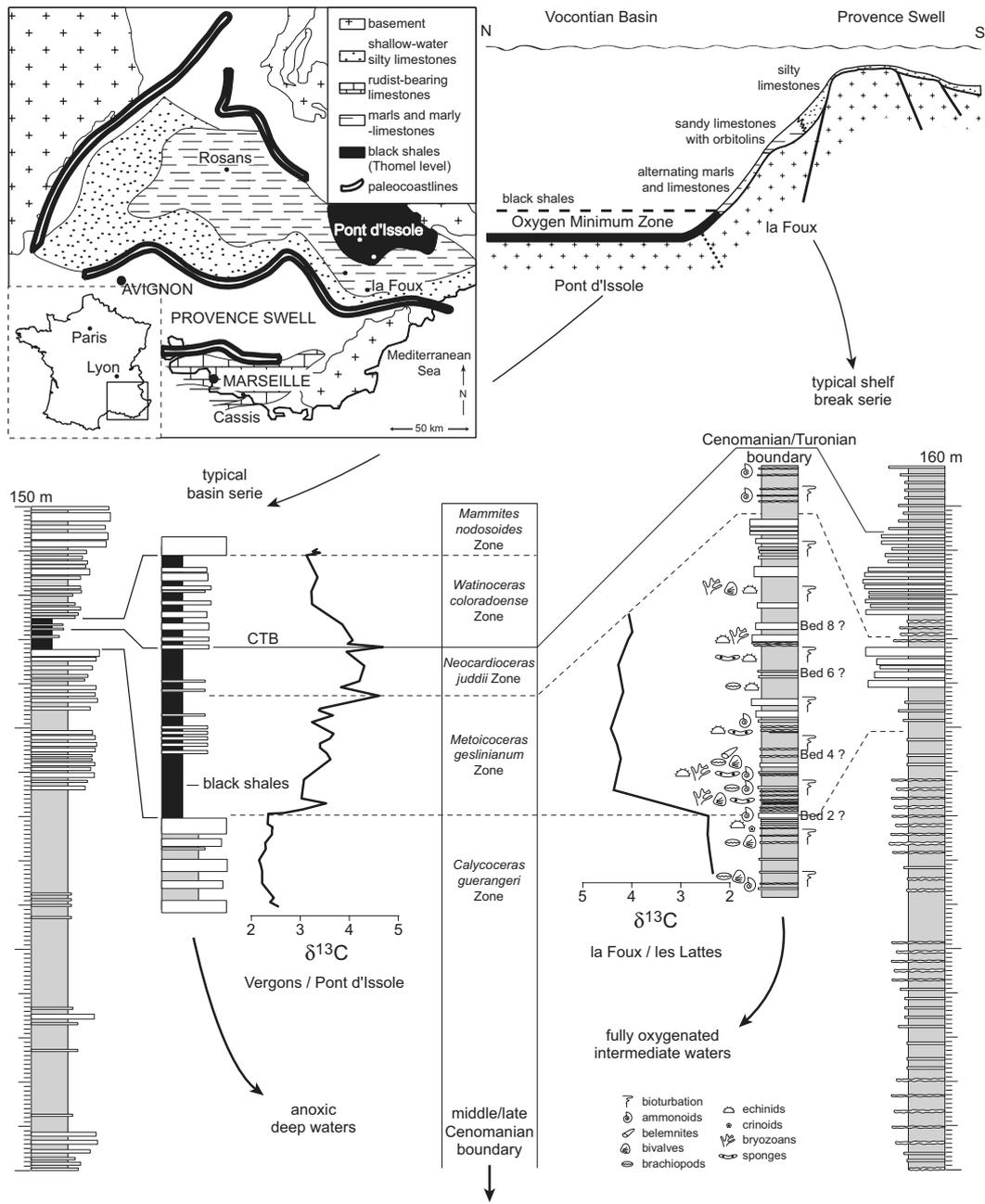


Fig. 2. Carbon isotope stratigraphy and ammonoid biostratigraphy of the la Foux and Vergons sections in the Vocontian Basin. $\delta^{13}\text{C}$ curve of la Foux after Gale & Christensen (1996). $\delta^{13}\text{C}$ curve of Vergons after Morel (1998).

comprises very thin (about one metre) variegated beds including a black shale (the Black Band), while in the southern province, the equivalent succession is up 13 m thick, devoid of black shales, and includes the Plenus Marls and basal Holywell Beds (base of the White Chalk Formation).

The extinction model of Jarvis et al. (1988), calling upon anoxia as a kill mechanism, is essentially based on data from the Dover section in the Anglo-Paris Basin. These authors argued that the extinction phase and the development of anoxic

waters occurred within the Plenus Marls. The Plenus Marls belong to the *Metoicoceras geslinianum* Zone (Wright & Kennedy 1981; Paul et al. 1999). The Black Band, whose correlation in the southern province is difficult, was initially considered the equivalent of the Plenus Marls (e.g. Rowe 1904; Hart & Bigg 1981; Hart & Leary 1989; Jeans et al. 1991; Hart et al. 1993). However, using carbon isotope data and macrofossil evidence, Wood & Mortimore (1995) and Morel (1998) independently demonstrated that the Black Band s.s. is younger

than the Plenus Marls and belongs to the *Neocardioceras juddii* Zone, whereas the Black Band s.l. started simultaneously with the Plenus Marls (i.e. *M. geslinianum* Zone). This correlation mirrors the German black shale facies successions, in which the main development of black shales overlies the Plenus Bed (i.e. *N. juddii* Zone) but also extends into the early Turonian (e.g. Hilbrecht & Hoefs 1986; Lehmann 1999).

The Cenomanian-early Turonian hemipelagic deposits of the Vocontian Basin, which locally may be up to 700 m thick, are composed of marly limestones alternating with silty marls (Fig. 2). In the latest Cenomanian these series are interrupted by several organic-rich layers (Thomel level), which vary in number from place to place (Morel 1998). These layers were deposited only in the deepest and most distal part of the basin (Crumière 1989; Crumière et al. 1990). In other basins, black shales documented were deposited in a similar depositional context (e.g. Tunisia, Robaszynski et al. 1994; Razgallah et al. 1994). In the Vocontian Basin, especially at Vergons, these layers have been intensively studied, including various methods (organic matter, Crumière 1989; Crumière et al. 1990; ichnofacies, Olivero & Gaillard 1996; foraminifers, Tronchetti & Groshény 1991, Groshény et al. 1992, Groshény & Tronchetti 1993; Groshény & Malartre 1997; ostracods, Babinot & Groshény 1993; palynology, Courtinat et al. 1991, Svobodová et al. 1998; carbon isotopes, Morel 1998). All these studies suggest that the Thomel level was deposited under anoxic, or at least hypoxic, conditions.

Only recently, the Thomel level has been precisely dated. Initially, it was assigned to the foraminifer zone *Whiteinella archeoretacea* (Groshény & Malartre 1997). Although the foraminifer zonation has a lower resolution than the ammonoid zonation, and the Vocontian sections with black shales contain only a few, non-age-diagnostic ammonoids, the Vocontian black shales coincide with the $\delta^{13}\text{C}$ positive excursion (Morel 1998). This major positive excursion is recognized worldwide and spans the latest Cenomanian *geslinianum* and *juddii* ammonoid zones.

Ammonoid biodiversity patterns

Monnet et al. (2003) analyzed the biodiversity patterns of ammonoids from the Anglo-Paris Basin and the Vocontian Basin. It should be noted that their study includes data from proximal to distal environments, thus lowering the preservation bias induced by the increasing sea level characterizing the late Cenomanian interval. Their study relies on the revision of the ammonoid zonation of these two basins, including numerous sections from each basin (see Monnet & Bucher 2002), thus lowering the sampling bias on documented ranges of ammonoids. Monnet et al. (2003) also used several statistical tests to assess the validity of the documented ammonoid biodiversity patterns.

For both basins, Monnet et al. (2003) demonstrated that during the early-middle Cenomanian, ammonoids show a high species richness characterized by a dynamic equilibrium. This

equilibrium rapidly changed at the middle/late Cenomanian boundary with a diversity drop to very low values of species richness (Fig. 3). It is also important to note that separate analyses of species richness for acanthoceratids and heteromorphs, the two essential components of the Cenomanian ammonoid community (41% and 23%, respectively), reveal that heteromorphs declined before the acanthoceratids. In the mentioned European basins, leiostraceous shells and most heteromorphs were already extinct at the middle/late Cenomanian boundary.

Discussion

The first issue is to compare diversity fluctuations of ammonoids with the timing of anoxic deposits. In both European basins, the biodiversity crisis of ammonoids started at the middle/late Cenomanian boundary (i.e. the base of the *Calyco-ceras guerangeri* Zone), while the spreading of anoxic waters onto the shelves started within the *geslinianum* Zone. Hence, ammonoid species richness declined one ammonoid zone before the spreading of anoxia (Fig. 3). This delay is significant, since the *Calyco-ceras guerangeri* Zone is supposed to last about 0.75 myr as established by radiometric ages from the Western Interior Basin (see Obradovitch 1993). Hence, the timing of ammonoid decline does not coincide with the development of anoxia, and anoxia alone cannot have triggered the decline of ammonoids. This decrease in ammonoid species richness may be seen as a result of a shift of facies belts caused by the late Cenomanian major flooding (see Smith et al. 2001). However, there is convincing evidence that the documented ammonoid decline is relatively unaffected by migration of facies or by preservation bias. They are documented for two basins with highly contrasted lithologies (e.g. chalks vs. limestone-marl alternations) and for each basin derived from sections ranging from proximal to distal depositional settings. Moreover, the differences in biodiversity patterns between acanthoceratids and heteromorphs suggest that there is no major preservation bias.

Most ammonoids are widely assumed to have had a nekto-benthonic mode of life, and thus should have gone extinct together with microbenthonic faunas with the onset of anoxic water episodes. Many studies have also related ammonoid shell morphotypes to environmental parameters, particularly to water depth (e.g. Ziegler 1967; Donovan 1985; Batt 1989; Westermann 1996). For the case of an upward migration of a poorly oxygenated water mass, Batt (1993) argued that highly ornamented, nekto-benthonic acanthoceratids should disappear first, while ancyloconic, pelagic heteromorphs should survive. From the north-west European record (see Monnet et al. 2003), the opposite pattern emerges. Hence, even if the time lag outlined above is ignored, ammonoid biodiversity patterns are not compatible with the development of anoxic waters.

These results raise the question of the influence of anoxia on the diversity of ammonoids. Since anoxic waters developed in the deepest parts of the basins, the question remains

whether anoxia did really spread onto the platforms and how it affected the biota. The model of Jarvis et al. (1988) assumed upward migration and spreading of anoxic waters onto platforms. However, in the Vocontian Basin, the black shales are clearly restricted to the deepest part of the basin (Fig. 2; Crumière 1989) and they are missing in slope and shelf environments. During the entire *geslinianum* Zone the sediments of outer shelf sections, such as at La Foux-Les Lattes, are intensively bioturbated and contain highly diverse benthonic faunas, including bivalves, sponges, irregular and regular echinoids, starfishes, bryozoans, inoceramids, as well as planktonic faunas (unpub. data, Fig. 2; see also Thomel 1992 and Gale & Christensen 1996). These faunas clearly developed under fully oxygenated conditions. It is noteworthy that the Thomel level consists of black shales alternating with bioturbated calcareous marls, which contain fish bones and intermediate-water planktonic foraminifers (Crumière 1989). This suggests that shallow environments of the Vocontian Basin were fully oxygenated. Hence, the model of a rising OMZ is not supported by data from the available sections in the Vocontian Basin, although the Thomel Level clearly indicates anoxic, or at least hypoxic, deep waters.

Initially Jarvis et al. (1988) based their model of anoxia-triggered extinction on data from the Dover section in the Anglo-Paris Basin. However, this model also takes into account data from other European basins. Although typical black shales are not present in the Dover section, Jarvis et al. (1988) argued that both sediments and microfaunas reflect the spread of oxygen-depleted environments during the Plenus Marls. However, Gale et al. (2000) argued that the decrease in diversity in microfauna, as well as changes in macrofauna, trace fossils and sedimentary geochemistry throughout the Plenus Marls interval were not related to decreasing oxygenation levels. As recalled by Smith et al. (2001), high-resolution studies are particularly sensitive to problems of stratigraphic completeness and sampling effort. Although numerous studies of various sections and basins claimed to document the progressive removal of shallower microfaunas due to the rising OMZ, the comparison of published data of various authors and from various sections of the Plenus Marls (e.g. Jarvis et al. 1988; Morel 1998; Paul et al. 1999; Keller et al. 2001) reveal that the extinction events of the various microfossil groups are rarely from the same bed. As pointed out by Gale et al. (2000) and Smith et al. (2001) the great late Cenomanian transgression may have induced local preservation biases generated by the spread of hemipelagic depositional environments onto the shelf. Local data are likely to be biased by local paleoenvironments and therefore may not reflect regional or global diversity patterns. To avoid such biases in studying the existence and the magnitude of a mass extinction, analyses should be performed at least at basin-scale in order to cover the whole array of environments (from coastline to deep basin). It is essential to note that studies of the global diversity of Cretaceous planktonic foraminifers (e.g. Hart 1999; Premoli Silva & Sliter 1999; Hart et al. 2002) have shown that anoxic events cannot be de-

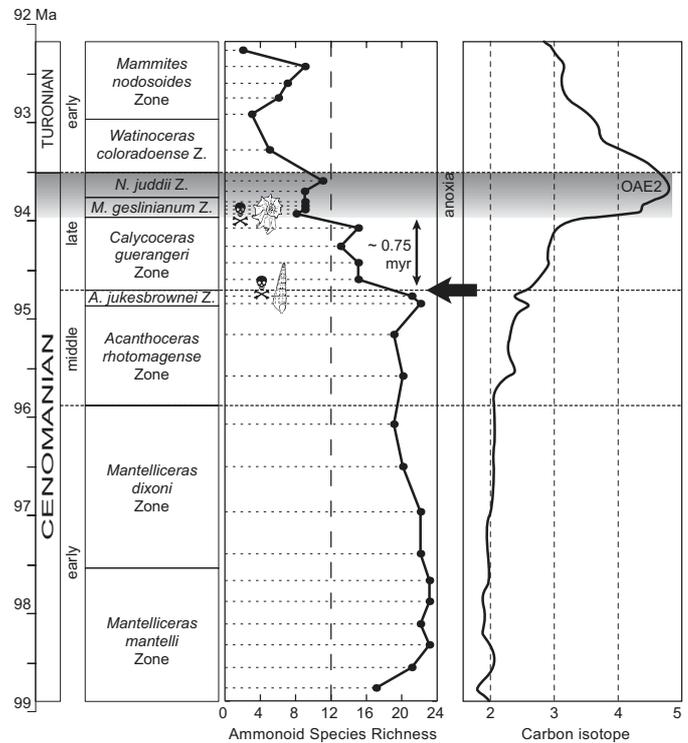


Fig. 3. Ammonoid species richness (after Monnet et al. 2003) and carbon isotope changes (after Voigt 2000) of northwest Europe during the Cenomanian-early Turonian. Note that ammonoid diversity dropdowns before the spreading of anoxia and that heteromorphs decline before acanthoceratids.

scribed as 'controlling factor' of the evolution of planktonic foraminifers. These studies emphasize that OAE2 is not associated with a drop in foraminifer diversity, but with a major turnover and a slight increase in taxonomic richness. According to Premoli Silva & Sliter (1999), the diversification of mesotrophic species (e.g. dicarinellids and whiteinellids) and the disappearance of specialized, oligotrophic species (e.g. rotaliporids) characterize the CTB turnover. These authors interpreted these trends as an indication of a change toward less stable and more eutrophic environments culminating in OAE2.

Finally, Corfield et al. (1990) studied stable oxygen and carbon isotopes of various morphotypes of benthonic and planktonic foraminifers from the Plenus Marls from Dover (Anglo-Paris Basin). As expected by the hypothesis of Jarvis et al. (1988), Corfield et al. (1990) found interesting evidence for vertical gradients of foraminifer habitats (e.g. *Rotalipora* in deeper water and *Hedbergella* in surface water). However, Corfield et al. (1990) also documented a very good correlation of planktonic and benthonic $\delta^{13}\text{C}$ records throughout the entire Plenus Marls interval, thus suggesting that during the late Cenomanian there was no change in the vertical $\delta^{13}\text{C}$ gradient. Consequently, subsurface oxygenation remained unchanged during the *geslinianum* Zone, and there is no clear relationship between the local extinction of benthonic and

planktonic foraminifers and macrofossil taxa and the expansion of the OMZ in the Anglo-Paris Basin. Hence, these available data do not support the model of a rising OMZ as major cause of the end-Cenomanian mass extinction.

Conclusions

The present study reveals and emphasizes several important results: (i) in both, the Anglo-Paris and Vocontian basins, the occurrence of oxygen-depleted water masses are confined to the deepest parts and limited in time to the *geslinianum* and *juddii* Zones (latest Cenomanian); (ii) in both basins shelf environments were fully oxygenated as demonstrated by oxygen and carbon isotope studies in the Anglo-Paris Basin and by abundant and highly diverse benthonic and planktonic faunas in the Vocontian Basin; (iii) in both basins the major changes of ammonoid faunas (diversity decline, evolutionary changes, taxonomic restructuring) started around the middle/late Cenomanian boundary (the *guerangeri* Zone), one ammonoid zone before the onset of the spreading of anoxic water masses in the deeper settings (the *geslinianum* Zone); and (iv) in both basins heteromorphs disappeared before acanthoceratids.

Although we agree that there is extensive evidence for widespread anoxia during the CTB interval in deep-water environments, our results rule out anoxia as the cause of the ammonoid crisis in northwest Europe. The latter predates by about 0.75 myr the evidence of anoxia in these basins. This lag is long enough to preclude anoxia as a killing-mechanism for ammonoids. Moreover, available data challenge the classic hypothesis that anoxia spread onto shelves during the late Cenomanian. Hence, even if OAE2 had contributed to the disruption of marine ecosystems and consequently to extinctions, other causative mechanisms must have initiated the moderate CTB mass extinction, at least for the ammonoid decline in north-west Europe. These results show the need to reassess the widely invoked causal link between anoxia and the moderate mass extinction at the Cenomanian/Turonian boundary.

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