

Bio- and chronostratigraphy of the Middle Triassic Reifling Formation of the westernmost Northern Calcareous Alps

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ABSTRACT

New finds of fossils including bivalves, ammonoids, brachiopods and palynomorphs from the Middle Triassic Reifling Formation significantly improve the age assignment for this unit in Liechtenstein and Vorarlberg. The lower part of the Reifling Formation is tentatively referred to the Late Anisian *Paracerasites trinodosus* Zone and somewhat older levels, whereas the uppermost part reaches the Ladinian *Protrachyceras archelaus* Zone (ammonoid zonation). The Middle Triassic successions of the study area are correlated with the coeval South Alpine reference section at Bagolino (Brescian Prealps), which also bears the Ladinian GSSP. The comparison shows that the Reifling Formation in the study area is age-equivalent with the South Alpine Prezzo Limestone and the Buchenstein Formation. A volcanoclastic layer in the upper part of the Reifling Formation at Flexenpass yields a U-Pb zircon age of 239.3 +/- 0.2 Ma. This value is slightly older than previously published minimum ages from equivalent horizons in the Southern Alps; the difference is thought to be mainly due to improved pre-treatment of zircons (annealing/chemical abrasion), which significantly reduces the effects of Pb loss. The new radio-isotope age further constrains the stratigraphical age of the Reifling Formation and supports the proposed biostratigraphy-based correlation of Middle Triassic successions in the Eastern and Southern Alps.

ZUSAMMENFASSUNG

Neue Funde fossiler Muscheln, Ammonoideen, Brachiopoden und Palynomorphen aus der mitteltriassischen Reiflinger Formation verbessern die Altersbestimmung dieser Einheit in Liechtenstein und Vorarlberg massgeblich. Der untere Teil der Formation wird provisorisch zur spätanisischen *Paracerasites trinodosus* Zone und älteren Horizonten gestellt, während der obere Teil bis in die ladinische *Protrachyceras archelaus* Zone (Ammonoideen-Zonierung) reicht. Die mitteltriassischen Sedimente des Untersuchungsgebietes werden mit dem Referenzprofil von Bagolino (Brescianer Voralpen) korreliert, welches den GSSP des Ladinian beinhaltet. Dieser Vergleich zeigt, dass die Reiflinger Formation des Untersuchungsgebietes gleich alt wie der Südalpine Prezzokalk und die Buchensteiner Formation ist. Die radiometrische U-Pb-Datierung von Zirkonen aus einer vulkanischen Aschenlage aus dem oberen Teil der Reiflinger Formation am Flexenpass ergibt ein Alter von 239.3 +/- 0.2 Ma. Dieser Wert ist älter als bekannte Minimum-Alter von entsprechenden Horizonten aus den Südalpen, und der Unterschied wird auf die verbesserte Zirkon-Vorbehandlung (Sintern gefolgt von chemischer Abrasion) zurückgeführt, wodurch sich die Effekte von Blei-Verlust massgeblich verringern lassen. Das neue Datum schränkt das stratigraphische Alter der Reiflinger Formation weiter ein und stützt die vorgeschlagene biostratigraphische Korrelation mitteltriassischer Sedimente der Ost- und Südalpen.

Introduction

The Middle Triassic successions of the Northern Calcareous Alps (NCA) part of the Eastern Alps predominantly consist of shallow marine to basinal carbonates with subordinate intercalations of fine-grained clastic sediments (Tollmann 1985). Platform carbonates and basinal deposits of Anisian to Carnian age occur in different tectonic nappe units but detailed reconstructions of the original platform-basin configuration are still controversial (see Ruffer & Bechstädter 1995 for a discussion of the situation in the western NCA).

For the heterogeneous basinal successions biostratigraphical information is as yet patchy and the resolution of published correlations of sections throughout the NCA is rarely better than at a substage level (e.g. Bechstädter & Mostler 1976). Reliable age control is limited to a small number of horizons bearing age diagnostic macrofossils in a few sections such as those around Grossreifling in the eastern NCA (Summesberger & Wagner 1972, Tatzreiter 2001 and references therein). Outcrops of pelagic successions including condensed Hallstatt-type limestones in the central NCA (e.g. Mandl 1984) produced Late Anisian to possibly Early Ladinian ammonoids reported

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long ago (Mojsisovics 1893; Diener 1900). In the southeastern NCA Hallstatt-type limestones alternating with an interval similar to the Reifling Formation also yielded age diagnostic conodonts and daonellas (Leithner & Krystyn 1984). From the Middle Triassic successions of the NCA north of the Inn Valley information on conodonts is available for a few coherent basinal sections (Bechstädt & Mostler 1974; Donofrio et al. 1979) as well as for isolated samples (Rüffer 1995). More recent and consistent conodont data, and a record of magnetic reversals, exist for a few sections of Late Anisian–Ladinian pelagic carbonates (Reifling Formation) in the eastern NCA (Gallet et al. 1998).

The Reifling Formation is a key stratigraphical interval of pelagic limestones with potential for significantly improved age calibration. In basinal settings at some distance from coeval carbonate platforms the stratigraphical thickness of this unit seems to show little variation and usually ranges in the order of a few tens of metres. Siliceous nodular limestones of the Reifling Formation represent a peculiar facies, which can hardly be distinguished from nodular limestones of the Buchenstein Formation in the Southern Alps. Both formations also bear discrete layers of acidic volcanoclastic debris, often with a distinct greenish colour. In the Southern Alps several of these so called *Pietra Verde* layers have been confirmed as primary volcanic ash beds and such tephra layers turned out to be ideal markers for precise correlation of sections up to 200 kilometres apart (Brack & Rieber 1993; Brack & Muttoni 2000). Moreover, the stratigraphical distribution of the volcanoclastic layers allowed a geochronological framework to be established on the basis of U-Pb single zircon ages (Mundil et al. 1996a). Therefore, if recorded with sufficient detail, the tuff beds in Middle Triassic successions of the NCA are expected to be similarly useful for correlation and radio-isotope dating.

Although the nature of volcanoclastic beds has long been recognised in the western NCA (Vidal 1953; Miller 1965; Sarnthein 1965, 1966; Hirsch 1966), stratigraphical sections with a published record of tuff layers suitable for layer tracing are rare (e.g. Bechstädt & Mostler 1974). Only for a restricted area in the westernmost NCA Kobel (1969) was able to convincingly demonstrate the lateral persistence of individual volcanoclastic beds. In the Reifling Formation east of the Inn Valley the thickness and frequency of such layers appear to be reduced. In the eastern NCA Gallet et al. (1998) used three principal tuff markers for the correlation of distinct Reifling intervals. However, in view of the larger number of tuff layers identified in similar successions around Grossreifling (Summesberger & Wagner 1972) the correlations suggested by Gallet et al. (1998) might be ambiguous.

Since the onset of Triassic research in the Alps the sedimentary units of the NCA have often been compared with the stratigraphically better constrained Middle Triassic sediments of the Southern Alps. This was based mainly on lithological similarities and the comparison of a few fossil levels (e.g. Arthaber 1906; Assereto 1971). More recently, correlations were also proposed for entire depositional sequences (Brand-

ner 1984; Rüffer & Zühlke 1995). However, a quantified assessment and comparison of the platform-basin evolution in the NCA is seriously hampered by the lack of sufficiently precise bio- and/or chronostratigraphical constraints.

In the Southern Alps the age constraints for Middle Triassic sediments have improved, especially in the past two decades. In order to achieve better resolved correlations similar progress is also required for data from the NCA. A re-assessment of available older (bio-)stratigraphical information could help in this context, but substantial progress will depend on the availability of accurately positioned additional fossils. In a step towards this goal we present and discuss new bio- and chronostratigraphical data from the Reifling Formation in the westernmost NCA, i.e. in Vorarlberg (Austria) and Liechtenstein. The new results are then compared with information from the Bagolino reference section in the Lombardian part of the Southern Alps. The study area was chosen because of its relatively uniform sedimentary successions including the basinal Reifling Formation for which essential stratigraphical information was readily available (Hirsch 1966; Kobel 1969). Moreover, recent palaeogeographical reconstructions (Brack et al. 1999) suggest that, in Triassic times, the western NCA were situated closer to the western (i.e. Lombardian) than to any other part of the Southern Alps.

Institutional abbreviation: PIMUZ, Paläontologisches Institut und Museum der Universität Zürich, Switzerland.

Setting, available age information and stratigraphy of the Reifling Formation in the study area

The Middle Triassic successions of Vorarlberg and Liechtenstein are relatively uniform, whereas the basinal successions in the NCA east of the Arlberg pass are often replaced by platform carbonates. The scarce diagnostic fossils provide only a coarse stratigraphical framework for these sediments. A no longer available single and presumably questionable specimen of the ammonoid “*Protrachyceras reitzi*”, as well as data on daonellas (*Daonella* cf. *indica*) and conodonts (*Gladigondolella tethydis*), suggest that the Anisian–Ladinian boundary is within the upper Reifling Formation (Trümpy 1916; Hirsch 1966; Kobel 1969). In analogy with the Southern Alps the occurrence of volcanoclastic layers in the upper Reifling Formation has also been considered an argument for an age of these beds close to the stage boundary.

For the present research stratigraphical successions have been studied in detail at the following sites (Fig. 1): Kulm near Triesenberg (Liechtenstein) (N47°07'11" E09°33'43" [WGS 84]), Nospitz near Malbun (Liechtenstein) (N47°05'34" E09°35'54"), Blankuskopf south of Nenzing (western Vorarlberg, Austria) (N47°06'01" E09°40'47"), and Flexenpass (eastern Vorarlberg, Austria) (N47°08'35" E10°10'01"). The mainly lithostratigraphical results of Hirsch (1966) and Kobel (1969) demonstrate that the pelagic Reifling Formation is bracketed by the shallow basinal carbonates of the Virgloria Formation

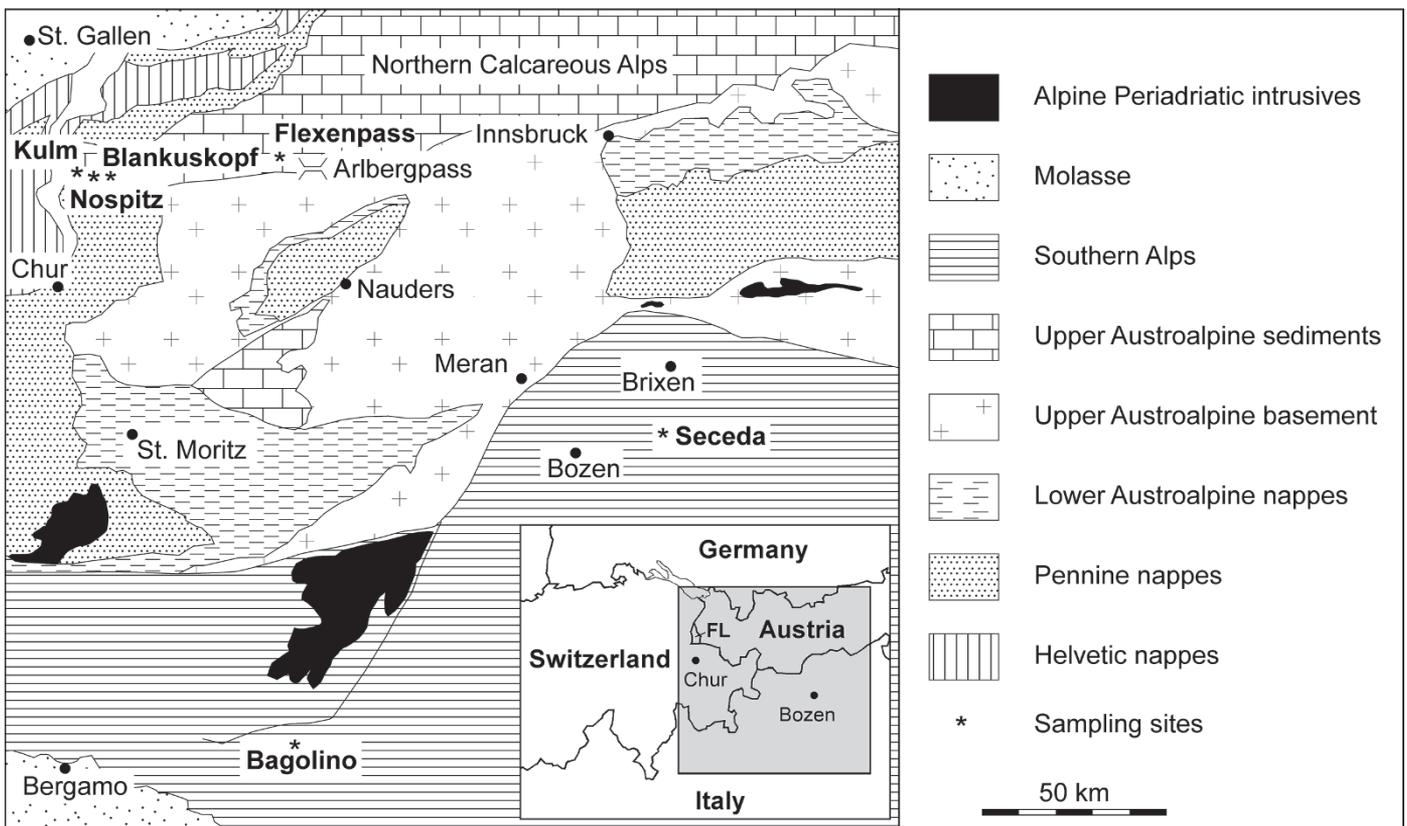


Fig. 1. Tectonic sketch map of eastern Switzerland, Liechtenstein, western Austria and northern Italy with the location of the studied sections. Tectonic map adapted from Gwinner (1971).

(below) and the dark shales and marls of the Partnach Formation (above). Here we adopt the stratigraphical denominations proposed by Bechstädt & Mostler (1974), who suggest using the name Virgloria instead of Gutenstein sensu Hirsch (1966) and Kobel (1969).

The Reifling Formation represents the upper part of the so-called “Alpine Muschelkalk” (see Kobel 1969 for a review and further sedimentological details) and it consists of nodular, micritic limestone beds, which, in their upper portion, show abundant chert in diffuse patches, bands and in distinct nodules. Following Hirsch (1966) and Bechstädt & Mostler (1974) the base of the Reifling Formation is defined at the upper boundary of a several metres thick horizon rich in crinoids [note that Kobel (1969) includes this level in his “Reifling Beds”]. Close to the top of the Reifling Formation a prominent 2–5 metres thick interval of Partnach-type shale is intercalated in the nodular limestones. These two distinct lithological markers bracket the Reifling Formation and can be recognised throughout Liechtenstein and Vorarlberg. In its upper half the formation also contains cm- to dm-thick volcanoclastic ash beds (*Pietra Verde*) and distinct bedding patterns of some of these layers allow an unambiguous correlation of the studied sections (Fig. 2).

New biostratigraphical data from the Reifling Formation and possible correlations with the Southern Alps

Macrofossils

In the studied sections macrofossils are generally very rare. Nevertheless, intense searching yielded a few bivalves, ammonoids and brachiopods (Plate I). Within the Blankuskopf section the brachiopod *Piarorhynchella trinodosi* is abundant in the lowermost part of the Reifling Formation. Only poorly preserved brachiopods were observed in the corresponding interval at Flexenpass. In the middle part of the Reifling Formation at Flexenpass a single specimen of *Daonella cf. longobardica* was found. More importantly, in all studied sections a horizon rich in the bivalve *Daonella tyrolensis* was observed in the upper part of the Reifling Formation, i.e. about 2 m below the laterally continuous, distinct interval of Partnach-type shales. At Kulm and Nospitz the shale interval also yielded some small ammonoids: several individuals of arcestids and a single specimen assigned to *?Arpadites* sp..

The brachiopod *Piarorhynchella trinodosi* typically occurs throughout the South Alpine Prezzo Limestone (Gaetani 1969). This unit corresponds to the *Paraceratites trinodosus*

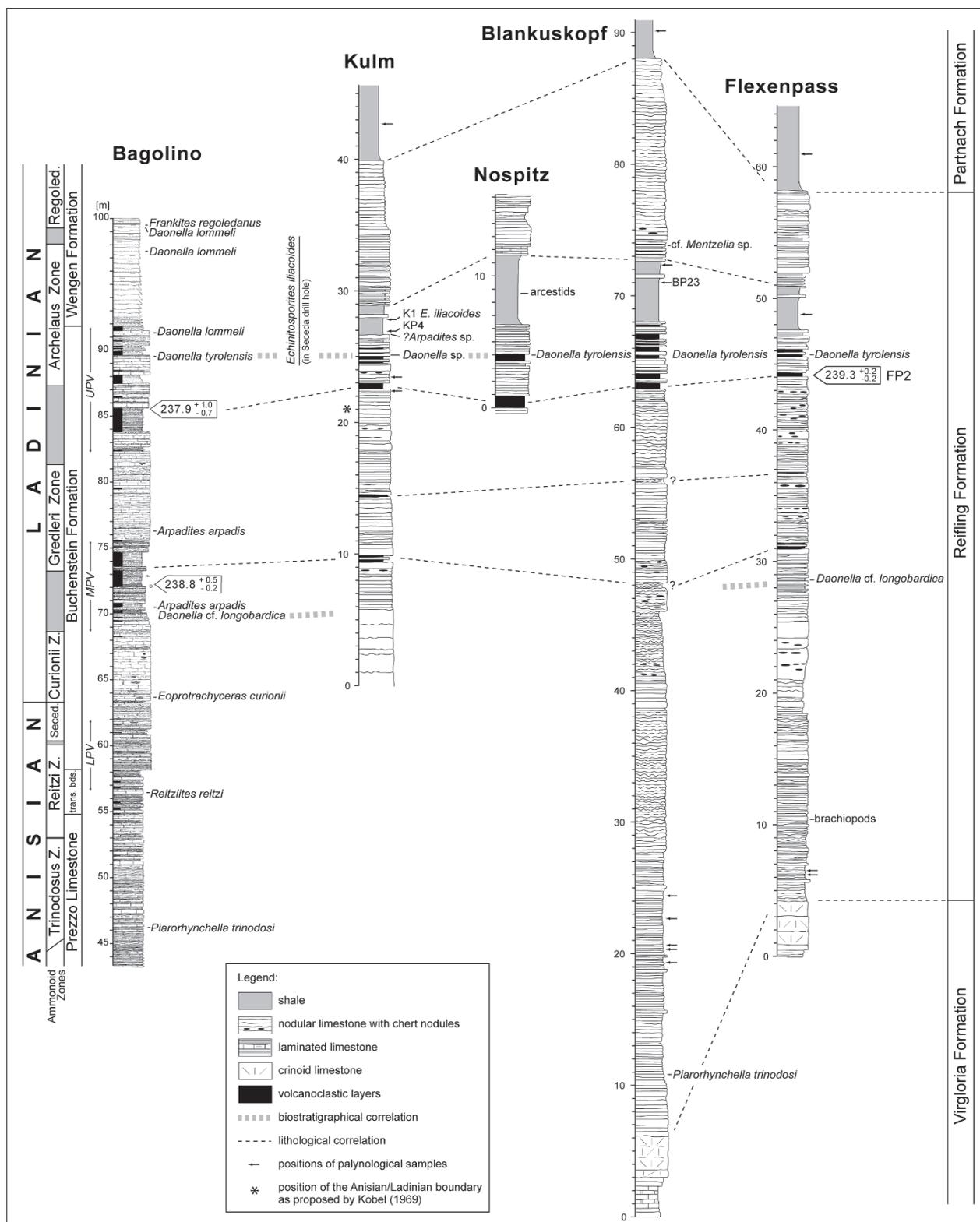


Fig. 2. Correlation of the studied sections of the Reifling Formation in the Northern Calcareous Alps (Liechtenstein, Western Austria) and their comparison with the South Alpine pelagic succession straddling the Ladinian GSSP at Bagolino (Northern Italy; for details see Brack et al. 2005; 0-m-level of Bagolino column corresponds to the base of the Prezzo Limestone). The main intervals with volcanoclastic layers in the Bagolino section are indicated as follows: LPV, MPV, UPV = Lower, Middle, Upper *Pietra Verde* respectively. Positions of seventeen palynological samples are indicated, six other analysed samples (yielding no palynomorph) come from the lower and middle part of the Partnach Formation at Blankuskopf. See text for a discussion of the radio-isotope ages.

Zone and older levels. The lower part of the Reifling Formation including the layers with *P. trinodosi* may thus correspond to this time interval.

Because of their global distribution and the presumably short ranges the thin-shelled bivalves of the subgenus *Daonella* have a high potential for biostratigraphical correlation of Middle Triassic sediments (Schatz 2001, 2004, 2005). Based on the stratigraphical distribution of different species of *Daonella*, Schatz suggested a *Daonella*-based biozonation for the Middle Triassic of the Southern Alps.

The species *Daonella longobardica* is best known from Ladinian platform slope carbonates in the western Southern Alps (Esino Limestone of the Bergamask Alps). In the Bagolino section specimens identified by Schatz (2005) as *D. cf. longobardica* occur in a distinct horizon of the middle Buchenstein Formation, i.e. at the transition of the *Eoprotrachyceras curionii* to the *Protrachyceras gredleri* ammonoid zones. These *Daonella* from Bagolino seem to be identical with a specimen from Flexenpass. The level with *D. cf. longobardica* at Flexenpass lies 2.5 m below a distinct thin volcanoclastic layer, which is here tentatively correlated with layers in the Middle *Pietra Verde* (MPV) interval at Bagolino (Fig. 2).

According to Schatz (2004) the distribution of *Daonella tyrolensis* is limited to the boundary interval between the *Daonella pichleri* Zone and the *Daonella lommeli* Zone, as defined in South Alpine sections (Bagolino, Seceda) and corresponding to the lowermost part of the *Protrachyceras archelaus* ammonoid Zone. Consequently the horizon with abundant *Daonella tyrolensis* in the upper part of the Reifling Formation is correlated with the uppermost Buchenstein Formation at Bagolino. This correlation is supported by the occurrence of the ammonoid *Arpadites* sp. in the shale interval close to the top of the Reifling Formation. In the Southern Alps ammonoids of this genus are thought to range from the *P. gredleri* to the *P. archelaus* zones (Brack & Rieber 1993).

A consequence of the new biostratigraphical results is the significant lowering of the position of the Anisian/Ladinian boundary when compared with Kobel (1969) who drew this limit above the 20m-level in the Kulm-section (Fig. 2). The stage boundary cannot yet be located precisely, but most likely falls in the interval below the layer with *Daonella cf. longobardica*. Moreover, the attribution to "*Protrachyceras reitzi*" of the lost ammonoid specimen collected by Trümpy (1916) cannot be confirmed in view of its stratigraphical position above the Partnach-type shale interval in the uppermost Reifling Formation. In the Southern Alps and in Hungary the species *Reitziites reitzi* is restricted to a narrow stratigraphical interval well below the Anisian/Ladinian boundary (cf. Bagolino column in Figure 2).

Palynology

Twenty three samples taken from different stratigraphical levels in the Reifling Formation and the Partnach Formation

from the Kulm, Blankuskopf and the Flexenpass sections (Fig. 2) were treated by standard HCl, HF and HNO₃ preparation methods and studied for their palynological content. Due to strong thermal alteration (see also Petschik 1989 and Ferreira Mählmann 1994) no palynomorph could be determined in samples from Flexenpass, whereas no significant alteration was observed in samples from Kulm and Blankuskopf. However, most samples from these localities show strongly degraded organic matter of variable quality of preservation. Only three samples from the shale interval in the upper part of the Reifling Formation (Fig. 2) provided palynomorphs suitable for determination (Plate II). Apart from pollen and spores these samples contain rich assemblages of acritarchs.

Sample KP 4 from Kulm (Fig. 2) yielded a few palynomorphs (*Calamospora* sp., *Baculatisporites* sp., *?Lycopodiacidites* sp.) as well as several acritarchs. The following palynomorphs were detected in sample K1 from the same section (Fig. 2, Plate II): *Aratrisporites scabratus* KLAUS 1960, *Aratrisporites cf. parvispinosus* LESCHIK 1955, *Ovalipollis* sp., *?Triadispora* sp., *Striatoabieites* sp., *Staurosaccites* sp., *Endosporites* sp., *Nevesisporites vallatus* DE JERSEY & PATEN 1964, *Microcachryidites* sp., *Verrucosisporites* sp., *Duplicisporites granulatus* (LESCHIK 1955) VAN DER EEM 1983, *Calamospora* sp., *Porcellispora longdonensis* (CLARKE 1965) SCHEURING 1970, *Gordonispora fossulata* (BALME 1970) VAN DER EEM 1983, *?Enzonasporites* sp., *?Cyclotriletes* sp., *?Infernopollenites* sp., *Baculatisporites* sp., *Deltoidospora* sp., *Concavisporites* sp., *Camerosporites secatus* LESCHIK 1955, *Echinitosporites iliacooides* SCHULZ & KRUTZSCH 1961, and several acritarchs. Sample BP 23 from Blankuskopf (Fig. 2) contains *Ovalipollis* sp., Circumpolles (*Partitisporites* sp., *Duplicisporites* sp., *Paracirculina* sp., *?Advisisporites* sp.), *?Triadispora* sp., *Aratrisporites* sp., *Nevesisporites vallatus*, *Todisporites* sp., *Cycadopites* sp., *Ephedripites* sp., *Deltoidospora* sp., and several acritarchs.

The stratigraphical distribution of spores and pollen in Middle Triassic successions of the Southern Alps and Hungary allowed Van der Eem (1983) and Brugman (1986) to establish discrete palynological phases. New results from the Seceda drill core (Brack et al. 2000; Hochuli & Roghi 2002) now provide an improved intercalibration of palynological data with ammonoid biostratigraphy. According to Van der Eem (1983) and Brugman (1986) the palynomorphs *Camerosporites secatus*, *Nevesisporites vallatus* and *Duplicisporites granulatus* first appear at the base of the *plurianulatus-secatus* phase. The base of this phase was originally correlated with the base of the *Eoprotrachyceras curionii* ammonoid Zone (Brugman 1986). However, new data from Seceda suggest that the *plurianulatus-secatus* phase more likely corresponds to the *Protrachyceras gredleri* ammonoid Zone (Hochuli & Roghi 2002). The species *Echinitosporites iliacooides* is an important stratigraphical marker with a range limited to the *secatus-dimorphus* phase of Van der Eem (1983). According to Brugman (1986), this phase corresponds to the *P. gredleri* and *P. archelaus* zones, but new data from the Seceda core show that *E. iliacooides* is

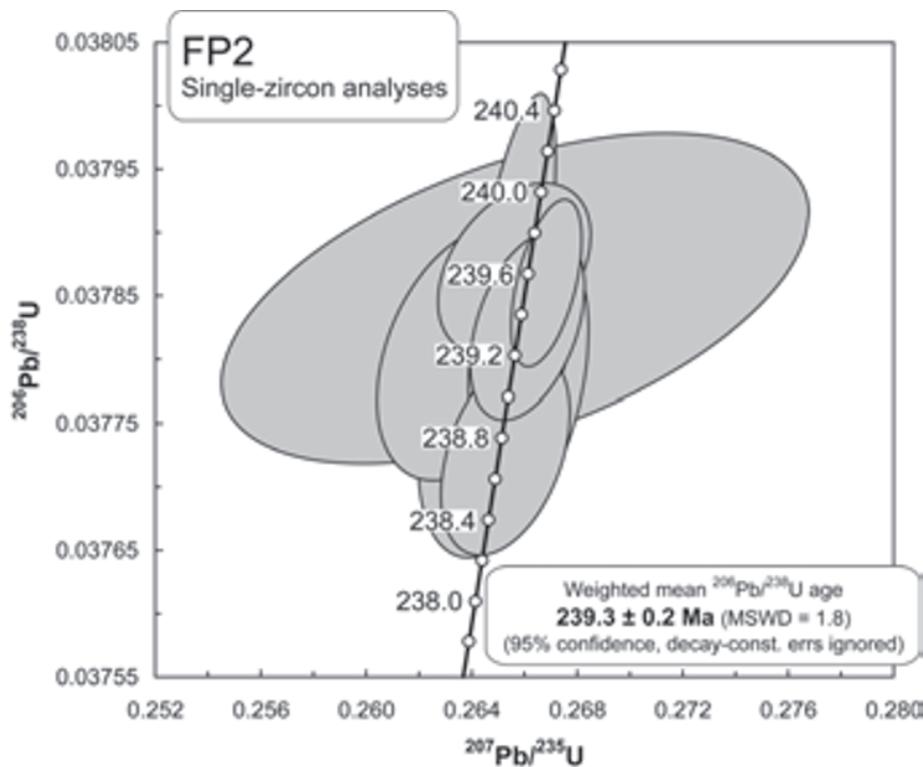


Fig. 3. Concordia diagram of analyses of zircons ($^{206}\text{Pb}/^{238}\text{U}$ vs. $^{207}\text{Pb}/^{235}\text{U}$) from the sample FP2 of the Flexenpass section (Reifling Formation, Northern Calcareous Alps, Austria). Data-point error ellipses are 2σ .

restricted to the *Protrachyceras archelaus* Zone (Hochuli & Roghi 2002). The shale interval close to the top of the Reifling Formation, yielding *Camerosporites secatus*, *Nevesisporites val-latus*, *Duplicisporites granulatus* and *Echinisporites iliacoides*, can be assigned to the *Protrachyceras archelaus* Zone supporting the suggested correlation of the uppermost Reifling and Buchenstein Formations.

Litho-biostratigraphical correlation of the studied sections with South Alpine successions

The basal Middle Triassic successions of eastern Lombardy and Giudicarie (Southern Alps) show the following stratigraphical order of formations (Assereto & Casati 1965, for a general stratigraphical scheme see also figure 2 in Brack & Rieber 1993): i) shallow marine Angolo Limestone, ii) pelagic Prezzo Limestone and Buchenstein Formation, and iii) clastic sediments of the Wengen Formation. The corresponding successions of the western NCA are similar in that the shallow-marine Virgloria Formation is followed by the deep-water Reifling Formation, which in turn is overlain by the clastic sediments of the Partnach Formation. In addition to this common overall trend of stratigraphical successions, some peculiar lithologies show striking similarities. Close analogies exist between the "Wurstelkalk"-intervals of the Virgloria (and Gutenstein) Formation and the strongly bioturbated and nodular intervals in the Angolo Limestone of eastern Lombardy. The top part of both units is often characterised by

crinoid beds and/or brachiopod-lumachella. The nodular lower Reifling Formation resembles the shale-poor nodular varieties of the Prezzo Limestone (e.g. in the Bagolino section) and the siliceous nodular limestones of the Reifling Formation are hard to distinguish from their counterparts in the Buchenstein Formation. Moreover, both units also contain volcanoclastic ash layers at different stratigraphical levels. The marls and shales of the Partnach Formation appear to be finer-grained when compared with the siliciclastic Lombardian Wengen Formation of latest Ladinian age. Lithologically similar to the Partnach shales are the local occurrences of the Lozio shales in the Southern Alps. The latter unit is restricted to exposures around the middle Val Camonica and situated stratigraphically above the Wengen Formation. Its base is already Carnian in age (Balini et al. 2000). Because no detailed biostratigraphical constraints are available for the Partnach Formation its correspondence with a distinct stratigraphical unit in the Southern Alps remains open.

In sections straddling the entire South Alpine Buchenstein Formation volcanoclastic layers are usually concentrated in three intervals: the Lower (LPV), the Middle (MPV) and the Upper *Pietra Verde* (UPV) (Brack & Muttoni 2000; Muttoni et al. 2004). Although volcanoclastic beds of the western NCA cannot be correlated on a layer basis, the two distinct intervals with *Pietra Verde* in the studied Reifling intervals (around the 31 m-level and between 44–46 m at Flexenpass) seem to correspond to the MPV and UPV, respectively (Fig. 2). This is in agreement with the biostratigraphical constraints.

The proposed correlation implies that the thickness values of corresponding intervals show little variation, not only within the westernmost NCA, but they also hardly differ from those of the Bagolino section, suggesting that the depositional setting and the sources of the pelagic sediments may have been uniform throughout large areas. On the basis of estimates made on Buchenstein sections in the Southern Alps (Brack & Muttoni 2000; Brack et al. 2007) the net rate of (not decompact) sediment accumulation of the nodular limestone facies of the Buchenstein Formation at Bagolino ranges between 5–10 m/Ma. The thickness of *Pietra Verde* layers is in most cases reduced in the NCA when compared with the South Alpine equivalents. This may be explained with a larger distance to the main centres of eruption.

U/Pb radio-isotopic dating of a tuff layer at Flexenpass: methods, results and comparison with data from the Southern Alps

Zircons from a volcanoclastic ash layer from the Flexenpass section (Fig. 2, 44.3 m level, horizon 12 in Hirsch 1966) were extracted by standard mineral separation techniques for U/Pb dating. The analyzed bed is about 30 cm thick and consists of well preserved, greenish material referred to as *Pietra Verde* (Hirsch 1966; Kobel 1969; Obenholzner 1991). The volcanoclastic layer shows distinct graded bedding with a few centimetres thick sandy layer at the base followed by siltstone.

Individual zircon grains were selected using transmitted light microscopy. Only euhedral, clear grains devoid of optically recognizable cores were chosen for analysis. All zircons were treated using the method of thermal annealing at 850°C for 36 hrs, followed by chemical abrasion with conc. HF in pressurised dissolution capsules at 220°C for 16 hrs in order to remove crystal portions affected by secondary Pb loss (Mundil et al. 2004; Mattinson 2005). Prior to dissolution the crystals were: (1) rinsed several times in conc. HNO₃, (2) cleaned in ultrasonically agitated aqua regia followed by several steps in clean HNO₃. Zircons were then transferred to perforated miniature PTFE capsules and spiked with ²⁰⁵Pb-²³³U-²³⁵U tracer solution. The capsules were placed on an elevated rack in a 125 ml digestion vessel, allowing vapour transfer dissolution in a mixture of 50% HF/conc. HNO₃ (30/1) at 220°C for 6 days. After dissolution the dried sample/tracer mixture was taken up in 20 µl 3 M HCl+5 µl 0.25 M H₃PO₄, dried down, and loaded together with silica gel+H₃PO₄ on out-gassed Re filaments. Isotope ratios were determined on a Micromass Sector 54 mass spectrometer using a Daly-type ion counter positioned behind a WARP filter. Pb (as Pb⁺) and U (as UO₂⁺) were run sequentially on the same filament. The ²⁰⁵Pb-²³³U-²³⁵U tracer solution was calibrated repeatedly against solutions derived from certified standards of isotopically pure ²⁰⁶Pb and natural U (NIST SRM-991 and CRM-145, respectively). To provide the highest level of confidence in the tracer calibration the standard solutions were themselves checked against

solutions prepared from other NIST-certified standards (NBLCRM111- A ²³³U, NBL-CRM135 ²³⁵U NIST SRM-981 and NIST SRM-982). All of these calibrations (effectively three independent calibrations each for U and for Pb) agreed within the precision of the measurements, so that we are confident of tracer Pb/U to within 0.15 %. In addition our tracer solution was checked against the similarly-calibrated tracer solution used at IGMR, ETH Zürich (itself checked against the tracer solution of the Royal Ontario Museum, as well as several other laboratories (Wiedenbeck et al. 1995). No inter-laboratory bias was detected by these cross-checks. In addition a cross check of the tracer (and entire procedure), involving the comparison of zircon U/Pb ages for the same rock (R33) between BGC and the ROM, showed agreement at the 1 permil level (Black et al. 2004) and most recently, an intercalibration study involving several laboratories, including the BGC, using zircons from R33 showed no significant inter-laboratory bias (say below the 2 permil level, Condon et al. 2005). Repeat measurements of the total procedural blank averaged 0.8 ± 0.3 pg Pb (U blanks were indistinguishable from zero), with ²⁰⁶Pb/²⁰⁴Pb = 18.55 ± 0.63, ²⁰⁷Pb/²⁰⁴Pb = 15.50 ± 0.55, ²⁰⁸Pb/²⁰⁴Pb = 38.07 ± 1.56 (all 2σ of population), and a ²⁰⁶Pb/²⁰⁴Pb-²⁰⁷Pb/²⁰⁴Pb correlation of +0.9. These ratios and uncertainties were propagated into the age and age-error calculations. Common Pb in excess of the analytical blank was assumed to have the same composition as given above, since using blank Pb or average 240 Ma crustal Pb has essentially no effect on the final result. Mass fractionation of U during analysis was controlled by the U double spike, whereas Pb mass fractionation was corrected by 0.15 ± 0.9 %/AMU (based on multiple analyses of NBS 981).

The U concentration of the analyzed zircons ranges between 200 and 660 ppm with one outlier containing 2200 ppm (Z03). Th/U varies from 0.27 to 0.39. Nine single zircons from a 5 kg rock sample were analysed yielding a weighted ²⁰⁶Pb/²³⁸U mean age of **239.3 ± 0.2 Ma** (MSWD 1.8; the uncertainty is augmented to ± 0.6 Ma if an additional 0.15% from tracer calibration is included for comparison with the ages in Mundil et al. (1996a) (Table 1 and Fig. 3). The U/Pb zircon ages of *Pietra Verde* horizons within the *P. gredleri* and *P. archelaus* zones at Bagolino and Seceda are slightly younger (BAG.06a, BAG.07, SEC.21 dated to 238.8 +.5/- .2 Ma, 237.9 +1.0/- .7 Ma, 238.0 +.4/- .7 Ma in Mundil et al. (1996a, b; see also Fig. 2). This is due to remaining effects of Pb loss, which has not been successfully removed by the then standard technique of air abrasion (Krogh 1982). Accounting for this complication the latter ages were deemed minimum ages in Mundil et al. (1996a). The recent re-analysis of a Latemar sample (Brack et al. 2007) shows that similar shifts towards slightly older ages may be expected also for values obtained on older levels in the Buchenstein Formation and at Latemar (Mundil et al. 1996a, 2003). Therefore we are confident that our new radio-isotopic age from Flexenpass is consistent with the bio- and lithostratigraphy-based correlation of the upper Reifling and the Buchenstein Formations.

Table 1. U-Pb isotopic data of zircons from the sample FP2 of the Flexenpass section (Reifling Formation, Northern Calcareous Alps, Austria).

^(a) Sample weight is calculated from crystal dimensions and is associated with as much as 50% uncertainty (estimated). ^(b) Total common Pb including analytical blank (analytical Pb blank is 0.8 ± 0.3 pg per analysis). ^(c) Present day Th/U ratio calculated from radiogenic $^{208}\text{Pb}/^{206}\text{Pb}$ and age. ^(d) Measured value corrected for tracer contribution and mass discrimination (0.15 ± 0.09 ‰/amu). ^(e) Ratios of radiogenic Pb versus U; data corrected for mass discrimination and fractionation, tracer contribution and common Pb contribution. ^(f) Correlation coefficient of radiogenic $^{207}\text{Pb}/^{235}\text{U}$ versus $^{206}\text{Pb}/^{238}\text{U}$.

Uncertainties of individual ratios and ages are given at the 2σ level and do not include decay constant errors. Ratios involving ^{206}Pb are corrected for initial disequilibrium in $^{230}\text{Th}/^{238}\text{U}$ adopting Th/U=4 for the crystallization environment. A weighted mean age of 239.3 ± 0.2 Ma (MSWD 1.8) is calculated from $^{206}\text{Pb}/^{238}\text{U}$ ages of individual analyses.

| Sample | $\mu\text{g}^{(a)}$ zirc. | ppm U | Pb _c ^(b) (pg) | Th ^(c) U | isotopic ratios | | | | | | isotopic ages | | | | |
|---------|------------------------------|-------|--|------------------------|--|--|-------|---|-------|---|---------------|--------------|---------------------------------------|---------------------------------------|--|
| | | | | | $^{206}\text{Pb}^{(d)}$ ^{204}Pb | $^{207}\text{Pb}^{(e)}$ ^{206}Pb | \pm | $^{207}\text{Pb}^{(e)}$ ^{235}U | \pm | $^{206}\text{Pb}^{(e)}$ ^{238}U | \pm | $\rho^{(f)}$ | ^{206}Pb ^{238}U | ^{207}Pb ^{235}U | ^{207}Pb ^{206}Pb |
| | | | | | | | % | | % | | | | | | |
| FP2.Z02 | 1,7 | 312 | 0,8 | 0,37 | 1637 | 0,05101 | 0,7 | 0,2653 | 0,7 | 0,03773 | 0,17 | 0,39 | 238.7 ± 0.4 | 239.0 ± 1.8 | 241.2 ± 16.0 |
| FP2.Z03 | 0,7 | 2184 | 1,1 | 0,39 | 3170 | 0,05086 | 0,4 | 0,2654 | 0,6 | 0,03785 | 0,35 | 0,68 | 239.5 ± 0.8 | 239.0 ± 1.3 | 234.6 ± 9.4 |
| FP2.Z04 | 0,4 | 284 | 0,8 | 0,37 | 348 | 0,05091 | 3,3 | 0,2657 | 3,4 | 0,03785 | 0,28 | 0,51 | 239.5 ± 0.7 | 239.2 ± 8.1 | 236.8 ± 75.2 |
| FP2.Z06 | 2,4 | 196 | 1,0 | 0,33 | 1156 | 0,05062 | 1,0 | 0,2639 | 1,0 | 0,03780 | 0,22 | 0,38 | 239.2 ± 0.5 | 237.8 ± 2.5 | 223.5 ± 22.6 |
| FP2.Z21 | 1,7 | 659 | 0,8 | 0,27 | 3513 | 0,05112 | 0,4 | 0,2669 | 0,4 | 0,03786 | 0,14 | 0,52 | 239.5 ± 0.3 | 240.2 ± 1.0 | 246.5 ± 8.2 |
| FP2.Z22 | 2,3 | 300 | 0,9 | 0,38 | 1835 | 0,05103 | 0,6 | 0,2662 | 0,7 | 0,03783 | 0,16 | 0,41 | 239.3 ± 0.4 | 239.6 ± 1.6 | 242.4 ± 14.1 |
| FP2.Z23 | 1,3 | 424 | 0,7 | 0,37 | 1870 | 0,05102 | 0,6 | 0,2660 | 0,6 | 0,03782 | 0,19 | 0,44 | 239.3 ± 0.5 | 239.5 ± 1.5 | 241.6 ± 13.4 |
| FP2.Z24 | 2,3 | 381 | 1,6 | 0,35 | 1304 | 0,05093 | 0,9 | 0,2652 | 1,0 | 0,03777 | 0,27 | 0,41 | 239.0 ± 0.7 | 238.9 ± 2.4 | 237.8 ± 21.1 |
| FP2.Z25 | 1,3 | 458 | 1,1 | 0,39 | 1366 | 0,05088 | 0,8 | 0,2657 | 0,9 | 0,03787 | 0,14 | 0,39 | 239.6 ± 0.3 | 239.2 ± 2.1 | 235.3 ± 19.4 |

Conclusions

New finds of macrofossils (*Daonella* cf. *longobardica*, *Daonella tyrolensis*, ?*Arpadites* sp., *Piarorhynchella trinosodosi*) and palynomorphs (*Echinosporites iliacooides*, *Camerosporites secatus*, *Nevesisporites vallatus*, *Duplicisporites granulatus*) significantly improve the age assignment of the Reifling Formation in Liechtenstein and Vorarlberg. The correlation of this unit with the interval of the Prezzo Limestone – Buchenstein Formation in the Ladinian GSSP section at Bagolino (Brack et al. 2005) is supported by biostratigraphical data and is consistent with radio-isotope age data. This age-equivalence is not unexpected since the Buchenstein and the Reifling Formations bear comparable lithologies and both units are parts of similar stratigraphical trends. The *Pietra Verde* levels recognised in the Reifling Formation of the study area correspond to layers of the Middle and Upper *Pietra Verde* intervals of the Buchenstein Formation of the Southern Alps. Finally, the Eastern Alpine Virgloria and the Gutenstein Formations are likely to be correlated with the Angolo Limestone of eastern Lombardy and Giudicarie, and the Partnach Formation of the western NCA might correspond to the Lombardian Wengen Formation. Elsewhere in the NCA the Reifling Formation also reaches higher levels (latest Ladinian or younger; Summesberger & Wagner, 1977; Gallet et al. 1998).

We are confident that additional bio/chronostratigraphical constraints can be generated also for Reifling successions to the east of Vorarlberg. Beyond an improved stratigraphical calibration such results will be essential for a quantified assessment of the platform-basin history of the NCA and its comparison with the evolution of coeval carbonate platforms in the Southern Alps.

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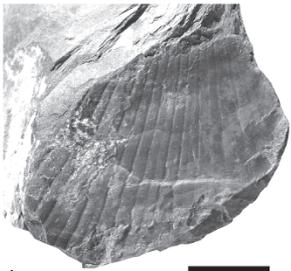
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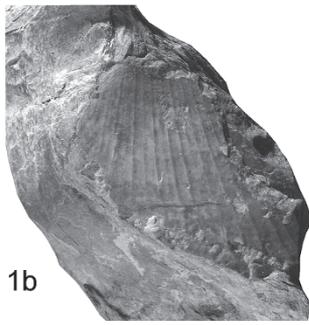
Plate I.

Macrofossils from the Reifling Formation of Liechtenstein (Kulm and Nospitz sections) and Austria (Blankuskopf and Flexenpass sections). Scale bars: 1 cm.

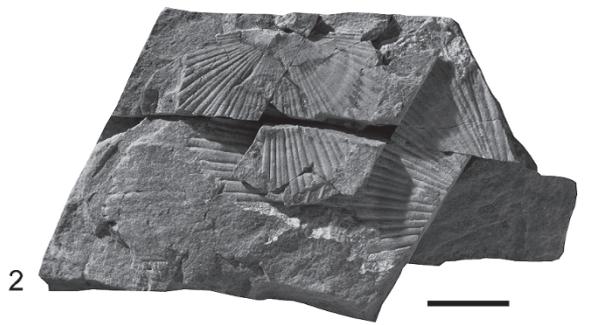
1) *Daonella* cf. *longobardica* KITTL 1912. Sinistral valve; a, b: positive, negative. Flexenpass, PIMUZ 24674. 2) *Daonella tyrolensis* MOJSISOVICS 1874. Dextral valve. Nospitz, PIMUZ 24670b. 3) *Daonella tyrolensis* MOJSISOVICS 1874. Dextral valve; a, b: negative, positive. Blankuskopf, PIMUZ 24676. 4) *Daonella tyrolensis* MOJSISOVICS 1874. Dextral valve. Nospitz, PIMUZ 24670a. 5) *Daonella tyrolensis* MOJSISOVICS 1874. ?Dextral valve. Flexenpass, PIMUZ 24675. 6) cf. *Mentzelia* sp. a, b: lateral, ventral views. Blankuskopf, PIMUZ 24673. 7) *Piarorhynchella trinodosi* (BITTNER 1890) DAGYS 1974. a, b, c: dorsal, lateral, anterior views. Blankuskopf, PIMUZ 24671. 8) ?*Arpadites* sp. a, b, c, d: sinistral, dextral, apertural, ventral views. Kulm, PIMUZ 24677.



1a



1b



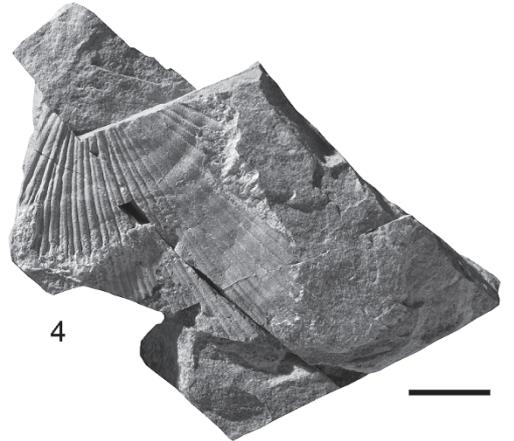
2



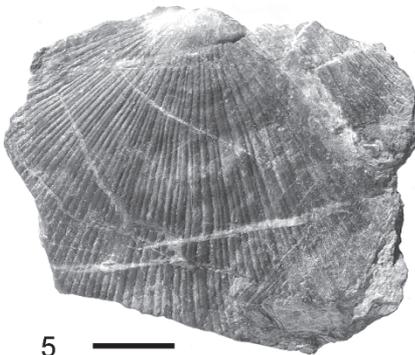
3a



3b



4



5



6a



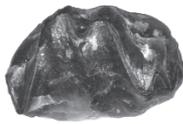
6b



7a



7b



7c



8a



8b



8c



8d



Plate II.

Palynomorphs from the Reifling Formation of Liechtenstein (Kulm section) and Austria (Blankuskopf section). Scale bars: 10µm (1–11, 13–17, 19, 20); 20µm (12, 18). Optical microscope images.

1a, b) *Echinitosporites iliacooides* SCHULZ & KRUTZSCH 1961, slide K1c. 2) *Partitisporites* sp., slide BP23. 3) *Camerosporites secatus* LESCHIK 1955, slide K1b. 4) *Baculatisporites* sp., slide K1c. 5) *Echinitosporites iliacooides* SCHULZ & KRUTZSCH 1961, slide K1b. 6) *Echinitosporites iliacooides* SCHULZ & KRUTZSCH 1961, slide K1. 7) *Duplicisporites granulatus* (LESCHIK 1955) VAN DER EEM 1983, slide K1b. 8) *Duplicisporites granulatus* (LESCHIK 1955) VAN DER EEM 1983, slide K1. 9) *Concavisporites* sp., slide K1. 10a, b) *Echinitosporites iliacooides* SCHULZ & KRUTZSCH 1961, slide K1c. 11) *Gordonispora fossulata* (BALME 1970) VAN DER EEM 1983, slide K1c. 12) *Porcellispora longdonensis* (CLARKE 1965) SCHEURING 1970, slide K1. 13) *Aratrisporites* cf. *parvispinosus* LESCHIK 1955, slide K1. 14) *Aratrisporites scabratus* KLAUS 1960, slide K1. 15) *Aratrisporites scabratus* KLAUS 1960, slide K1c. 16) *Calamospora* sp., slide K1. 17) *Ovalipollis* sp., slide K1. 18) *Ovalipollis* sp., slide K1b. 19) *Striatoabieites* sp., slide K1b. 20) *Nevesisporites vallatus* DE JERSEY & PATEN 1964, slide K1.

